

Couplage : Interactions océan-atmosphère au sein des cyclones tropicaux

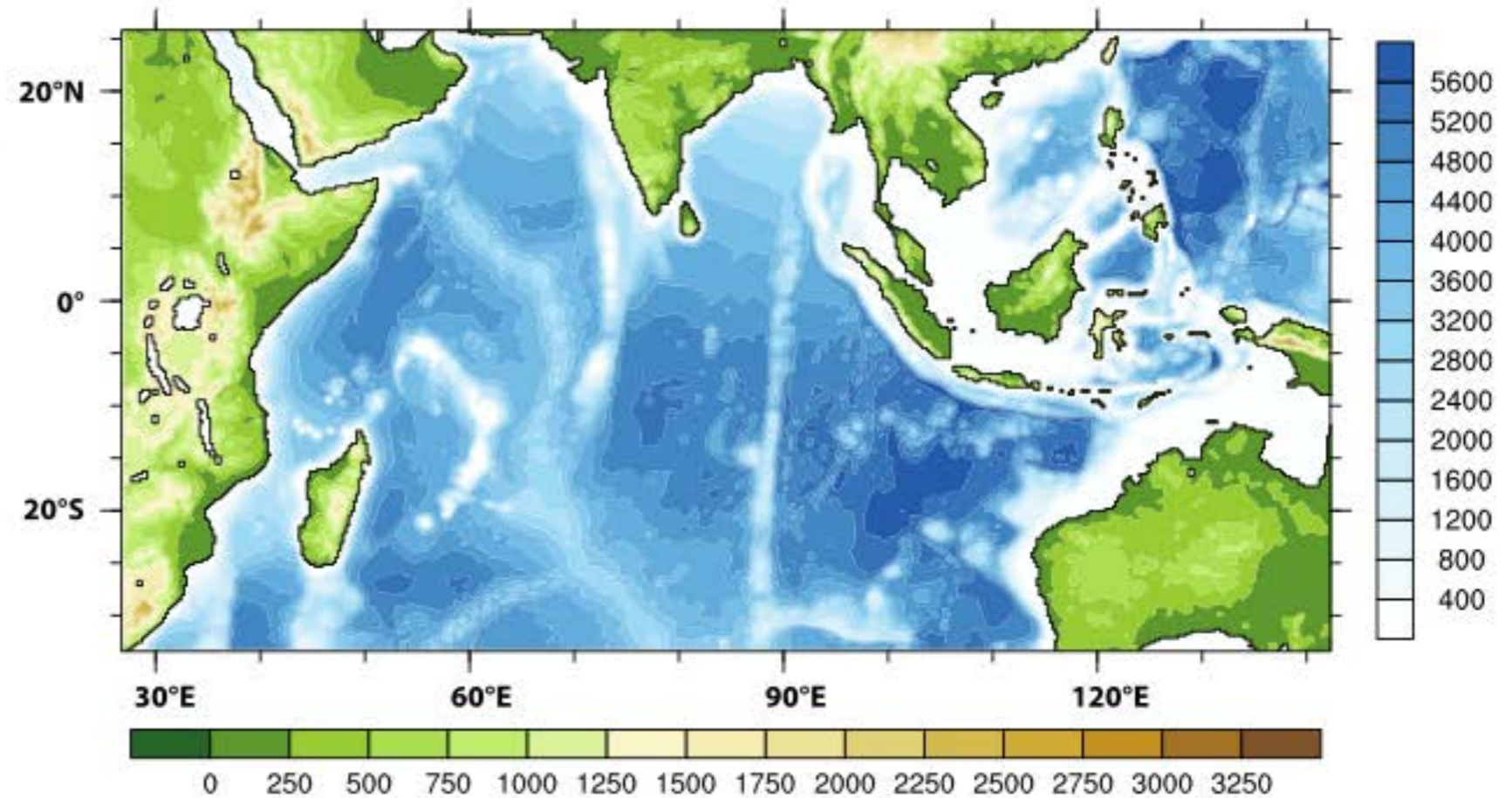
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Menkes, S. Jullien, S. Neetu

Impact of convection scheme

KF - BMJ

Configuration

NEMO-WRF 0.25°
1990 - 2009



Samson et al. 2014

- WRF single-moment six-class microphysics scheme WSM6 [Hong and Lim, 2006]
- Goddard shortwave radiation scheme [Chou and Suarez, 1999]
- Rapid Radiation Transfer Model (RRTM) for longwave radiation [Mlawer et al., 1997]
- Yonsei University planetary boundary layer YSU [Noh et al., 2003]
- Monin-Obukhov surface layer parameterization, and the four-layer Noah land surface model [Chen et al., 1996]
- drag parameterization over the ocean is based on the work of Donelan et al. [2004]

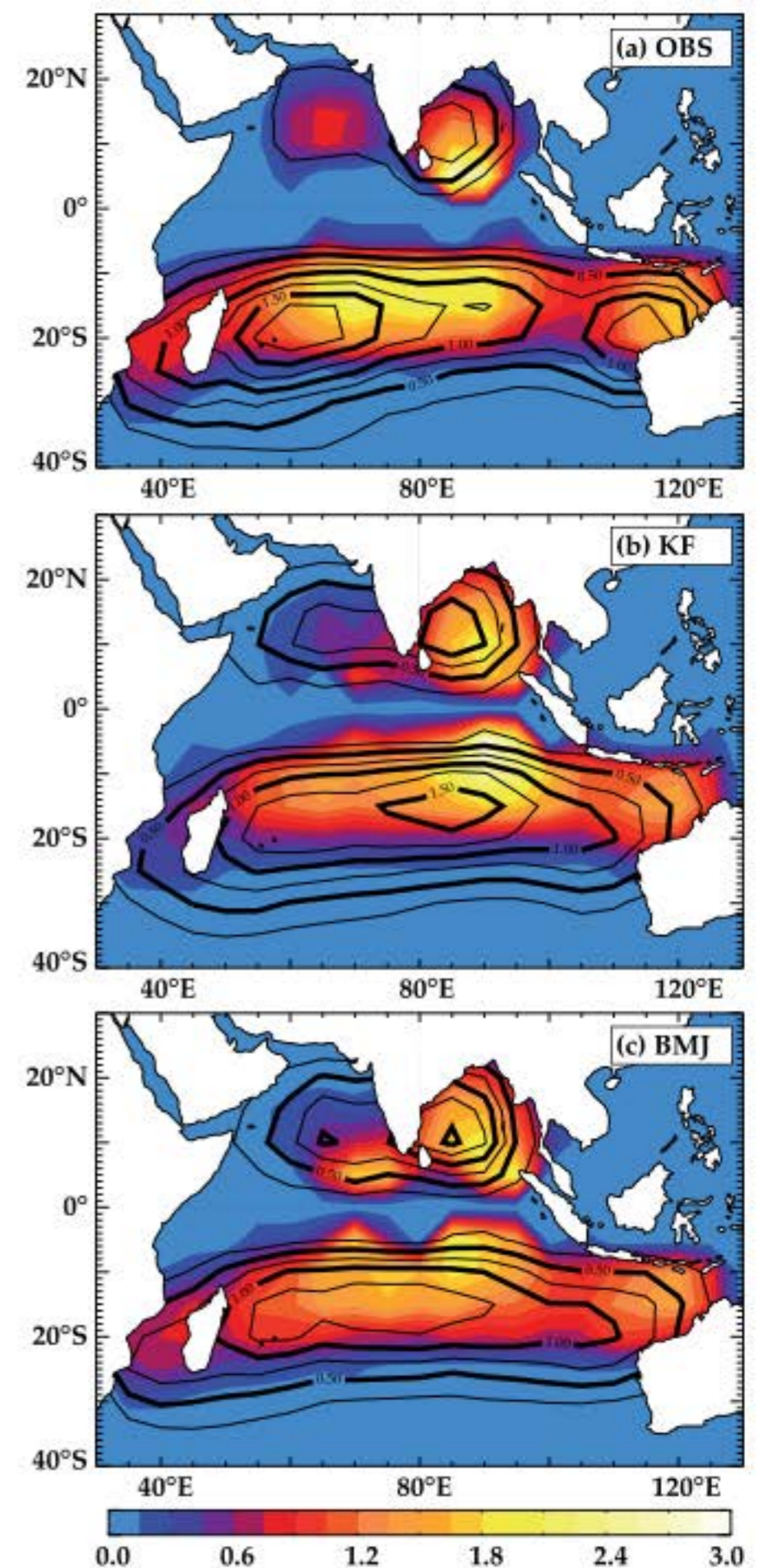
Cyclogenesis

Betts-Miller-Janjic (BMJ) scheme [Janjic', 1994]

- moist convective adjustment scheme,
- thermodynamic profile is adjusted toward quasi-equilibrium state.
- Activation mainly determined by thermodynamics (CAPE)

Kain-Fritsch (KF) scheme [Kain, 2004]

- simulate a vertical rearrangement of mass that eliminates CAPE
- triggered based on the grid-resolved vertical motion



Cyclogenesis seasonal cycle

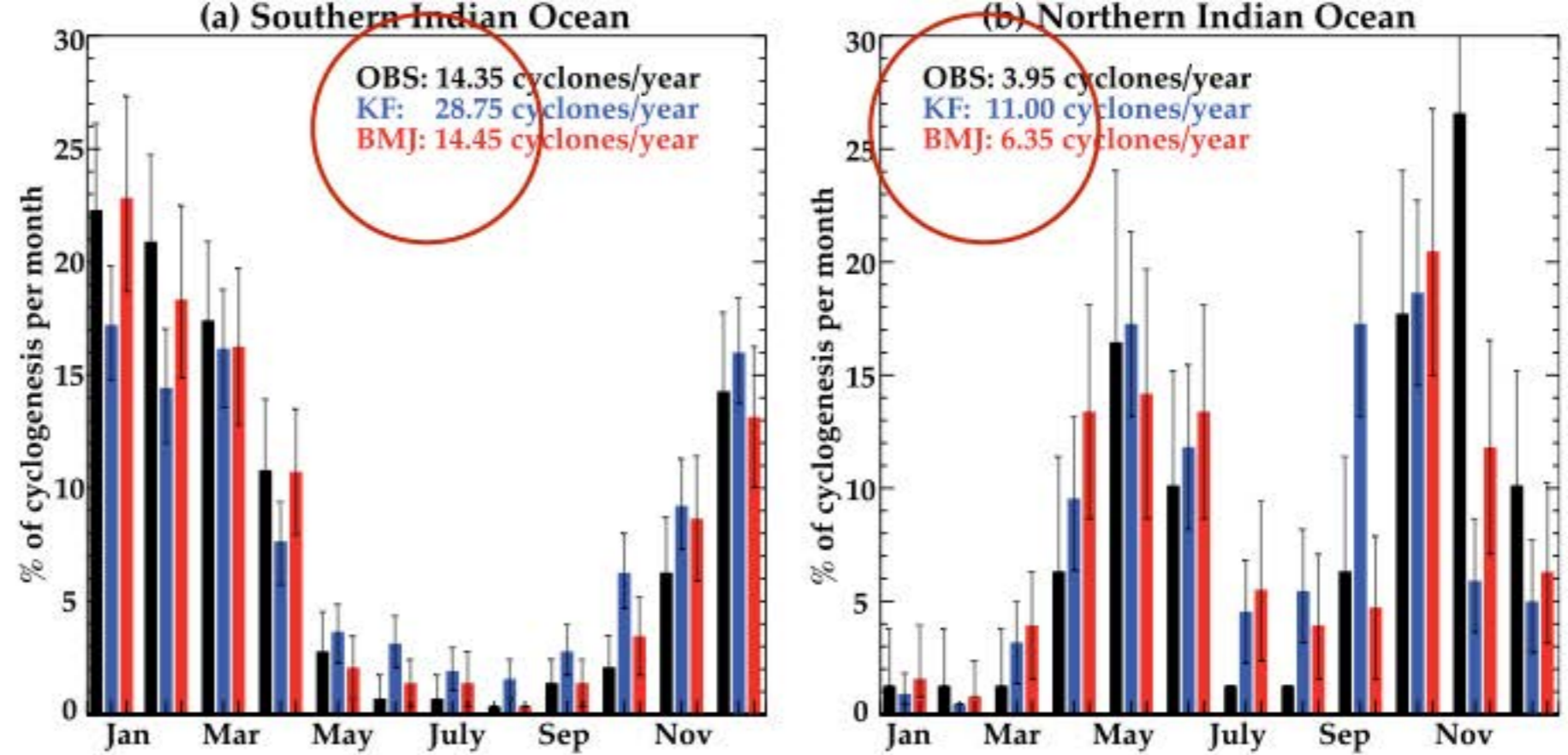


Figure 11. Seasonal cycle of relative cyclogenesis (in % of the total number of cyclones per month) for observations (black), KF (blue), and BMJ (red) experiments for (left) the southern and (right) northern Indian Ocean. The whiskers display 95% confidence interval on monthly cyclogenesis values based on a Student's *t* test.

Intensity distribution

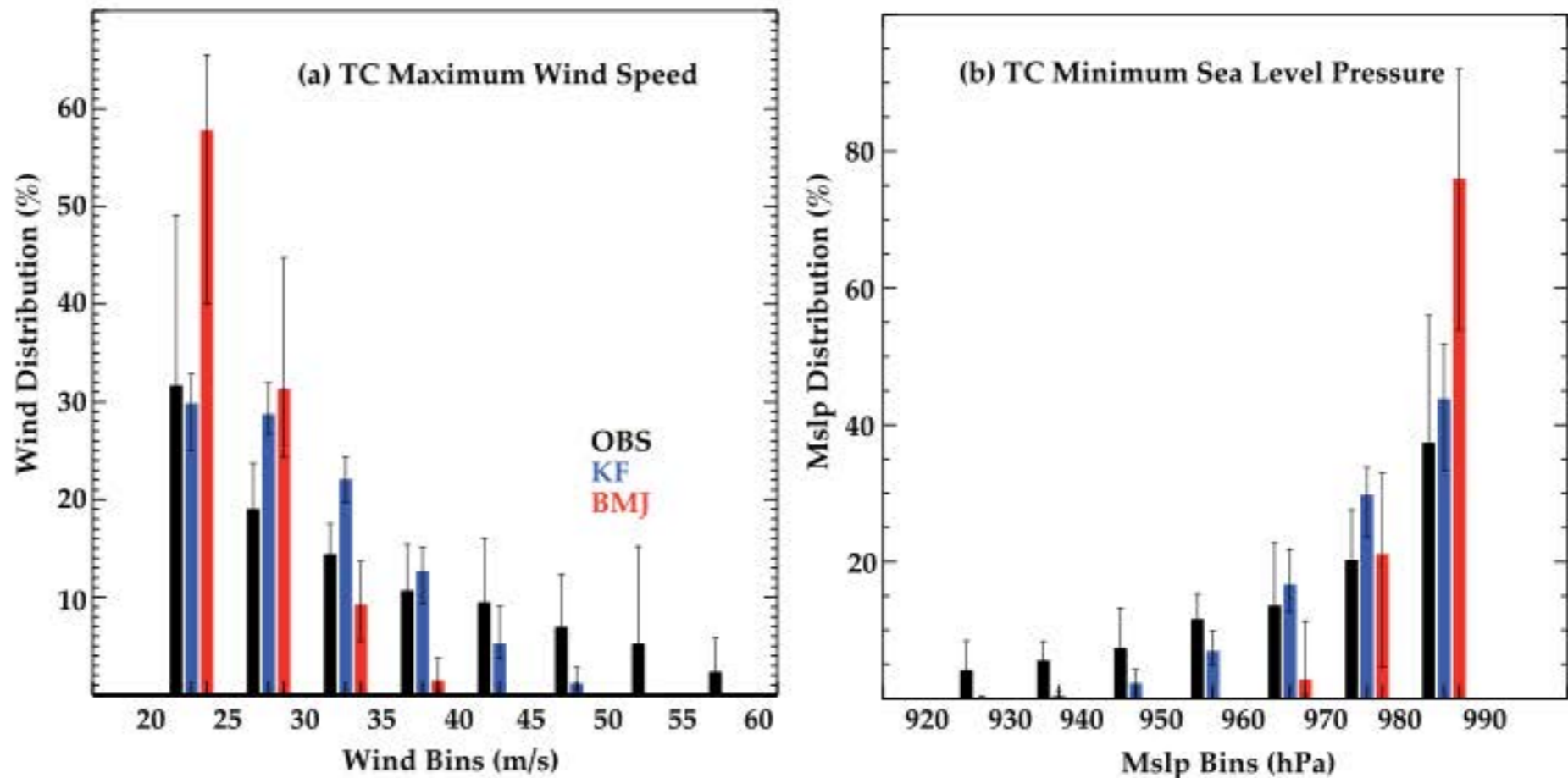
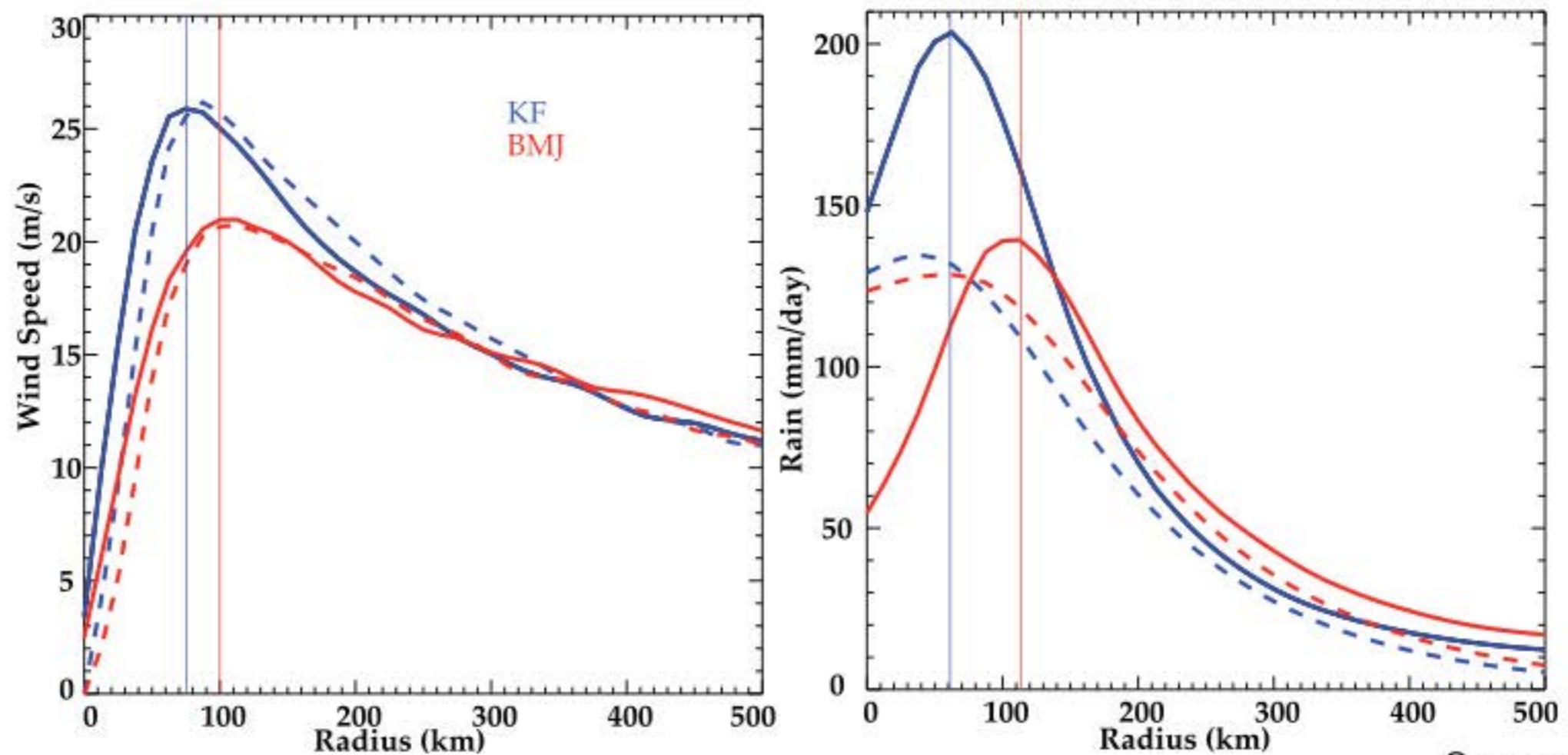


Figure 13. (left) Maximum wind intensity and (right) minimum sea level pressure distribution in observations (black), KF (red), and BMJ (blue) simulations for all Indian Ocean TCs.

radial profile



Samson et al. 2014

Figure 14. (a) Composite of TCs radial wind profile as a function of the distance to the TC center directly derived from the model (plain) and reconstructed using *Willoughby et al.* [2006] formulation (dashed) for KF (blue) and BMJ (red) simulations. (b) Composite of TCs rainfall profile directly derived from the model (plain) and reconstructed using *Tuleya et al.* [2007] formulation (dashed) for KF (blue) and BMJ (red) simulations. Vertical lines indicate the wind speed and precipitation maxima for both simulations.

=> one example of the sensitivity to one parameterization
many others...

Impact of coupling

Coupled - Forced

methodology

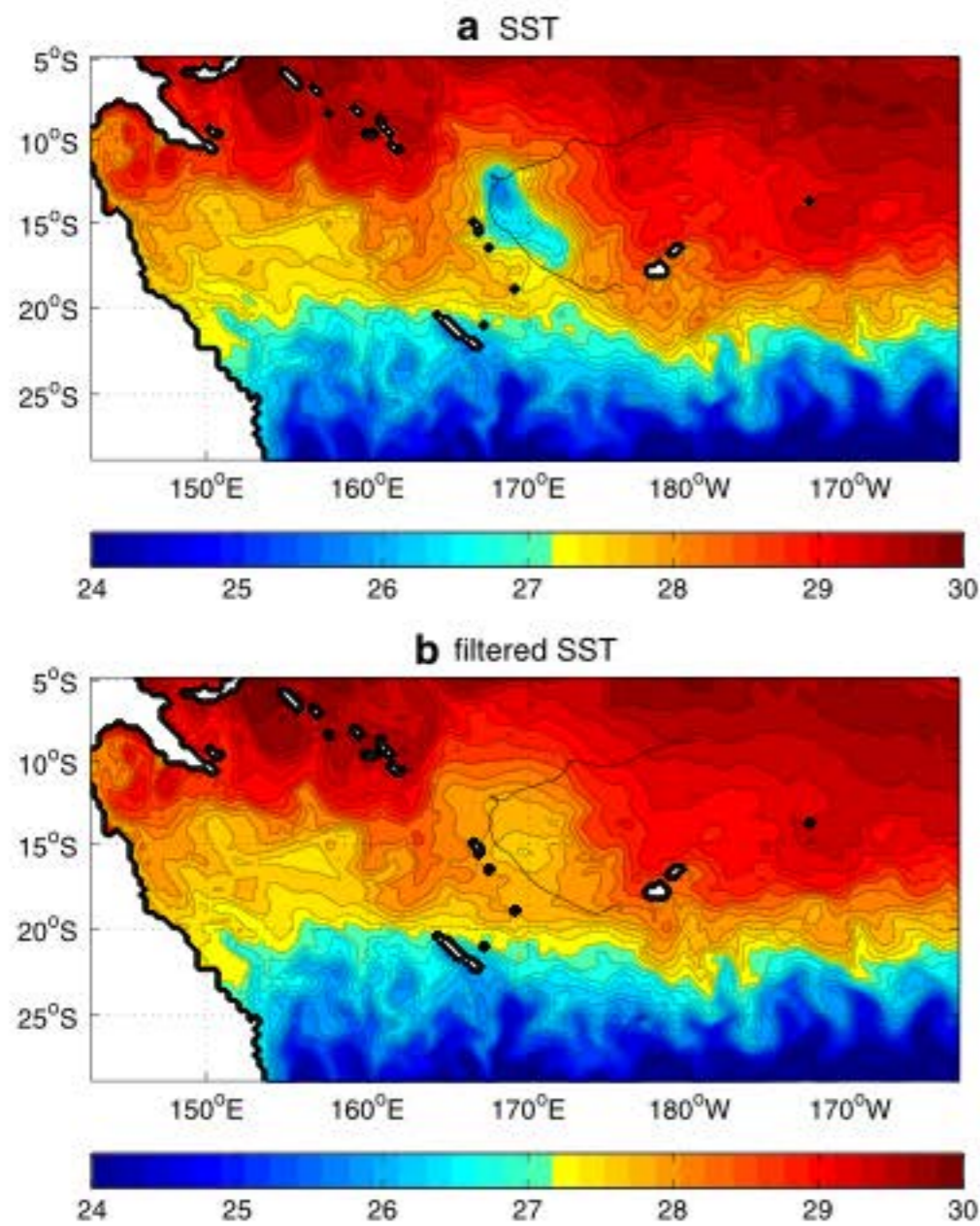


Fig. 1 Example of a TC cold wake event and its filtering procedure: **a** SST field from the coupled model ($^{\circ}\text{C}$), **b** filtered SST field used in the forced model WRF-ROMS $1/3^{\circ}$

Jullien et al. 2014

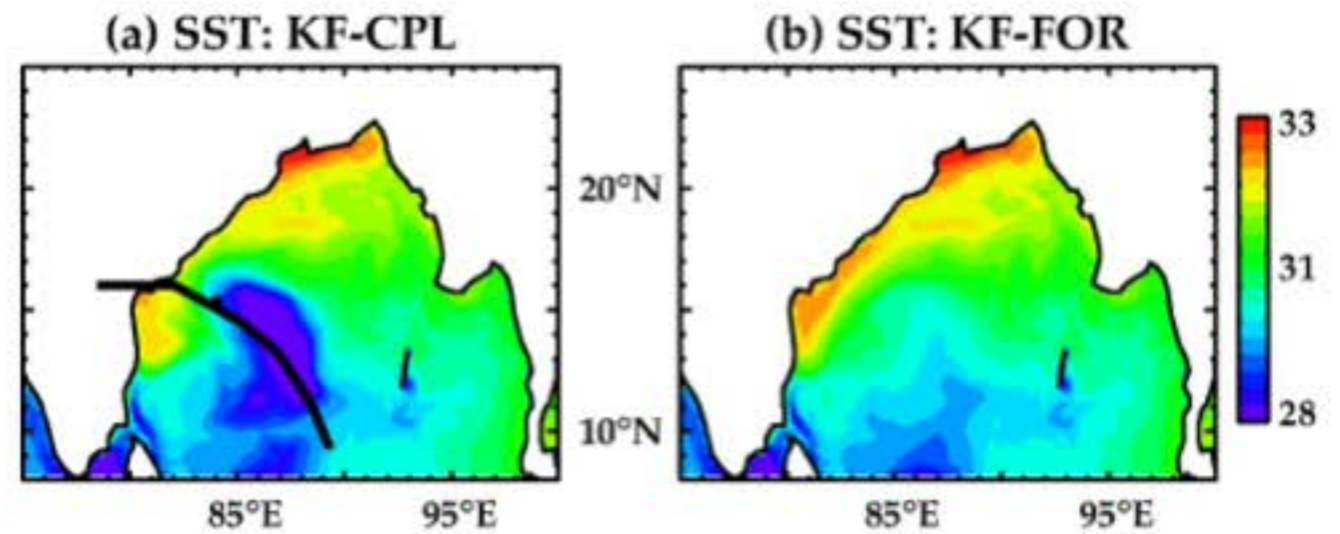
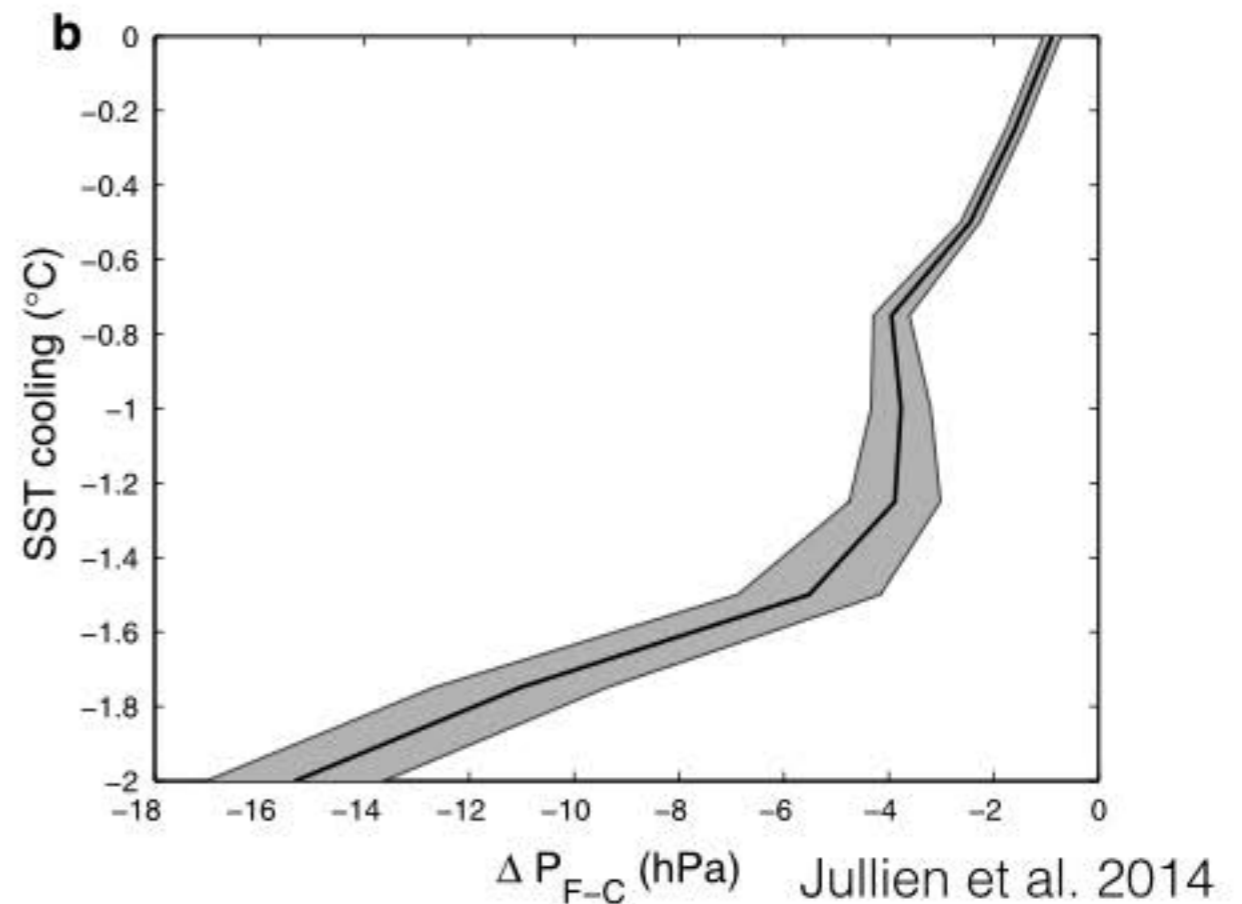


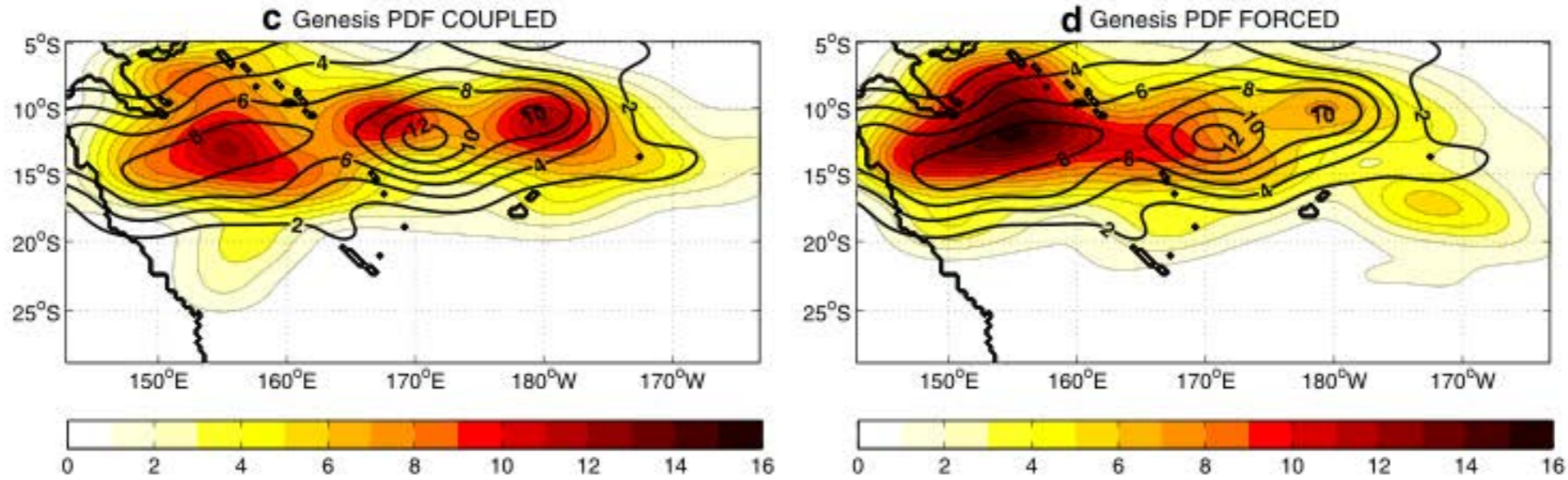
Figure 2: Sea surface temperature ($^{\circ}\text{C}$) snapshot for the **(a)** KF-CPL and **(b)** KF-FOR simulations. The SST boundary condition of the KF-FOR simulation is obtained after filtering out (see text for details) the TC cold wake from the SST field shown in **(a)**. The KF-CPL simulation TC track from the is shown by a black line on panel **(a)**.

Lengaigne et al. 2018



ΔP_{F-C} (hPa) Jullien et al. 2014

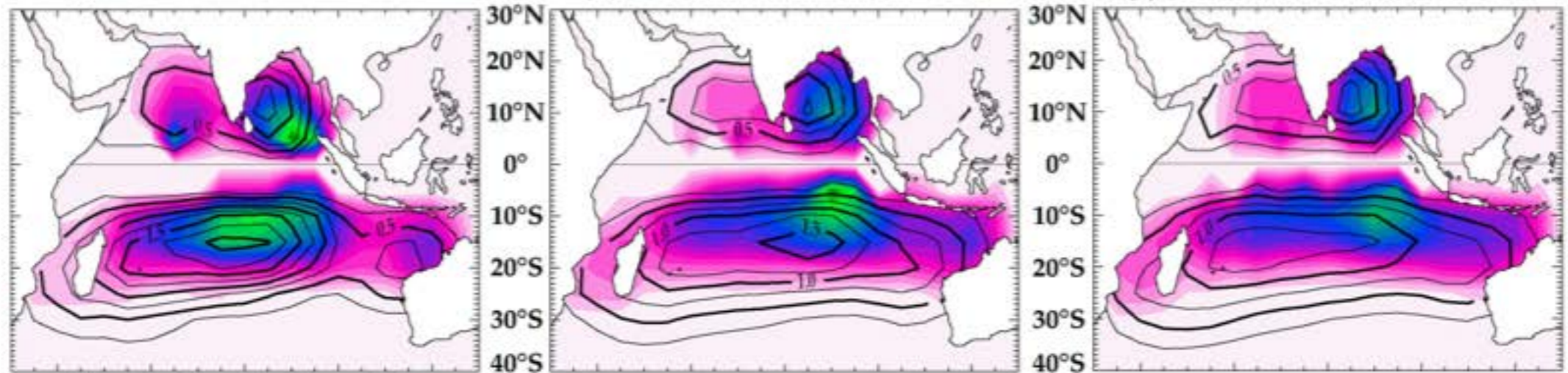
Cyclogenesis



(a) TCs distribution: OBS

(b) TCs distribution: KF-CPL

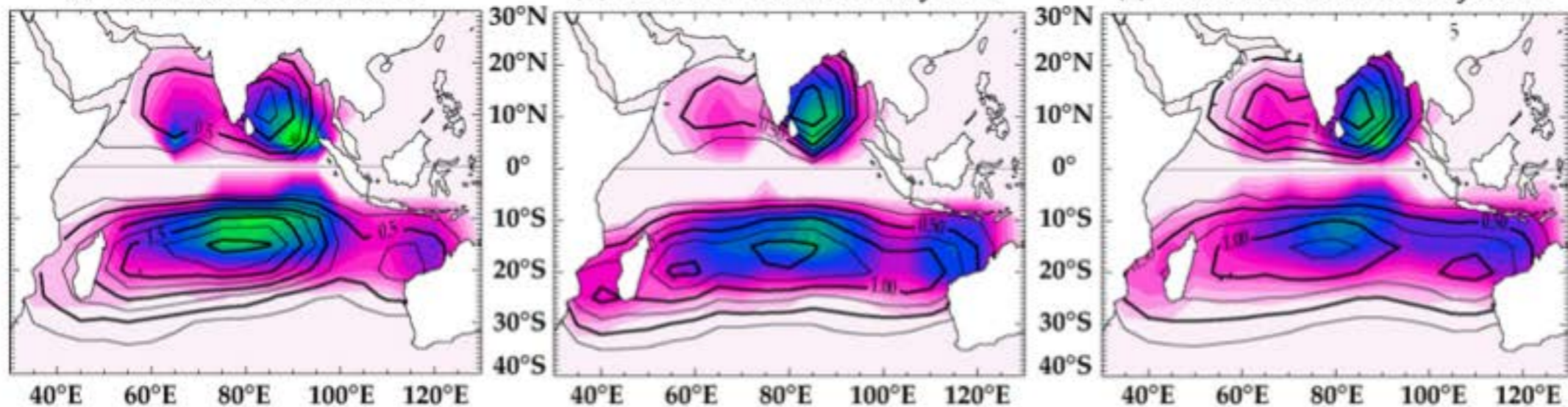
(c) TCs distribution: KF-FOR



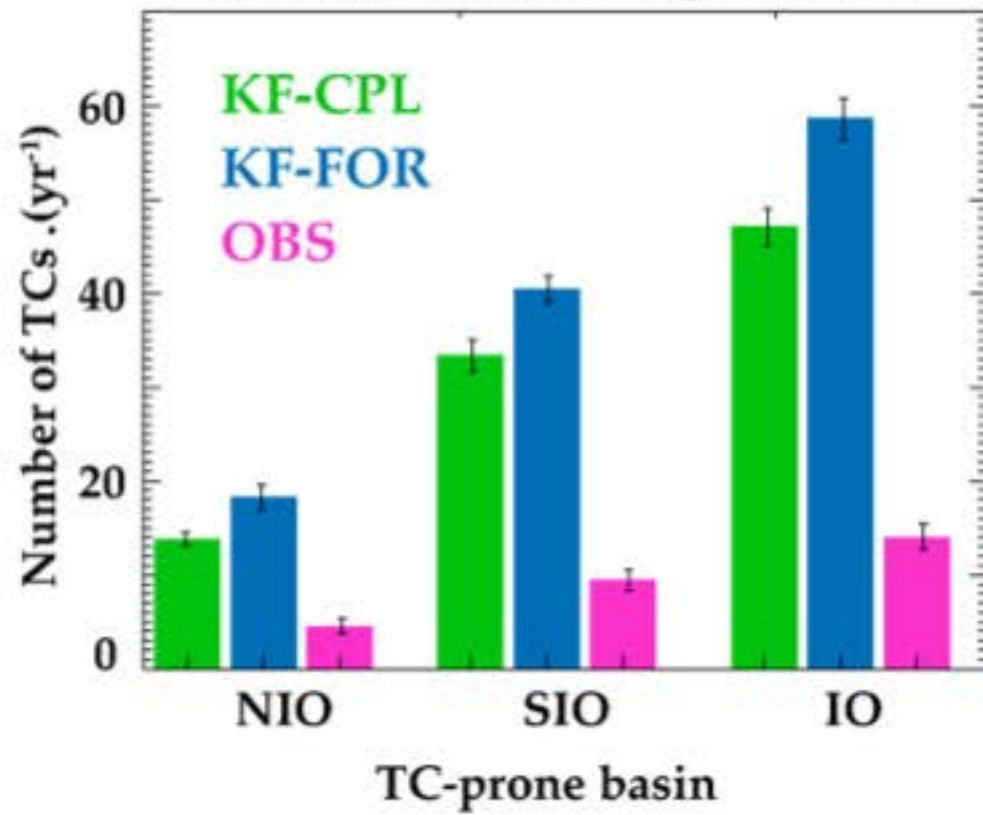
(a) TCs distribution: OBS

(b) TCs distribution: BMJ-CPL

(c) TCs distribution: BMJ-FOR



(d) Number of TCs per basin



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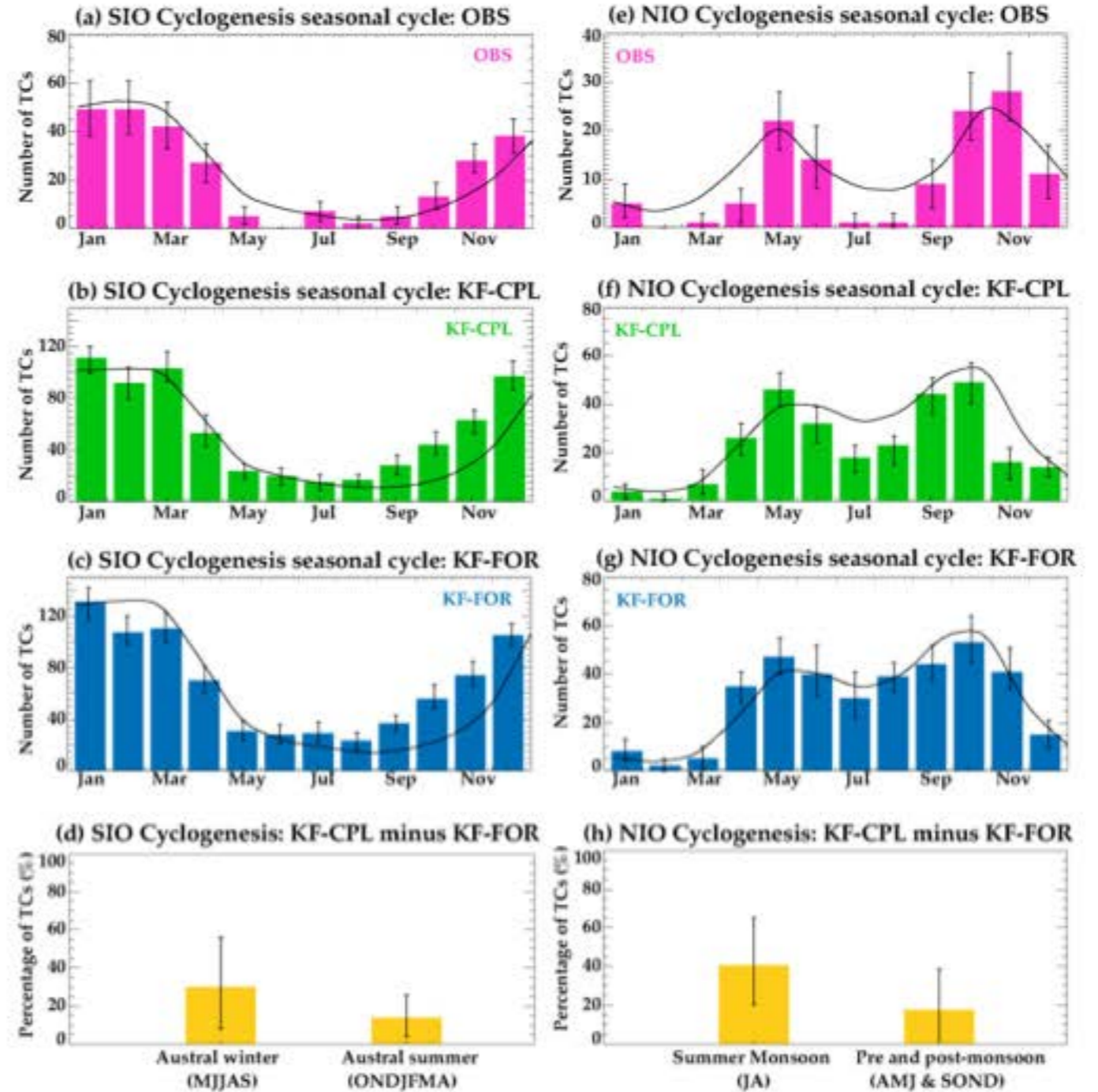
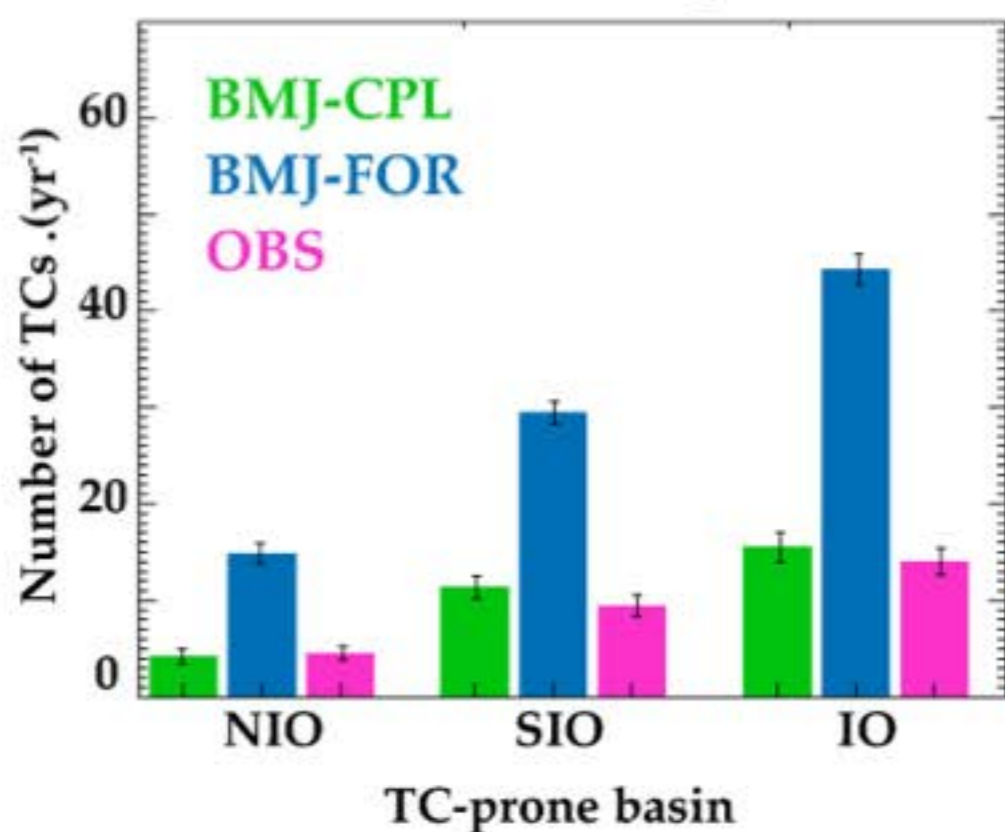


Figure 3: Histogram of the percentage of TCs occurring each calendar month in the (left column) northern and (right column) southern IO for (a,e) observations, (b,f) KF-CPL and (c,g) KF-FOR simulations. The monthly climatological evolution of the corresponding GPI index (northern IO: 40°E-100°E; 0°-25°N and southern IO: 30°E-130°E; 0°-25°S) is overlaid as a black curve. Reduction of the number of TCs (%) in KF-CPL relative to KF-FOR for the cyclonic and non-cyclonic seasons for (d) northern and (h) southern IO. On all panels, the whiskers display the 90% confidence interval based on a student t-test.

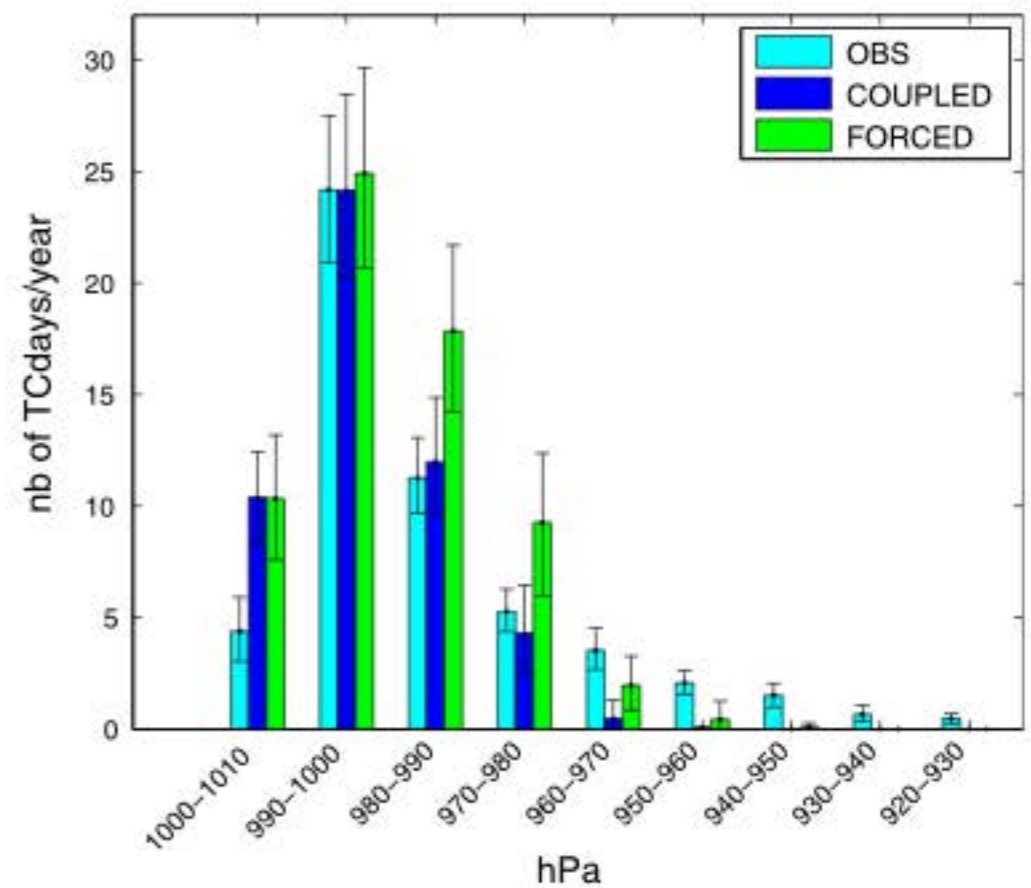
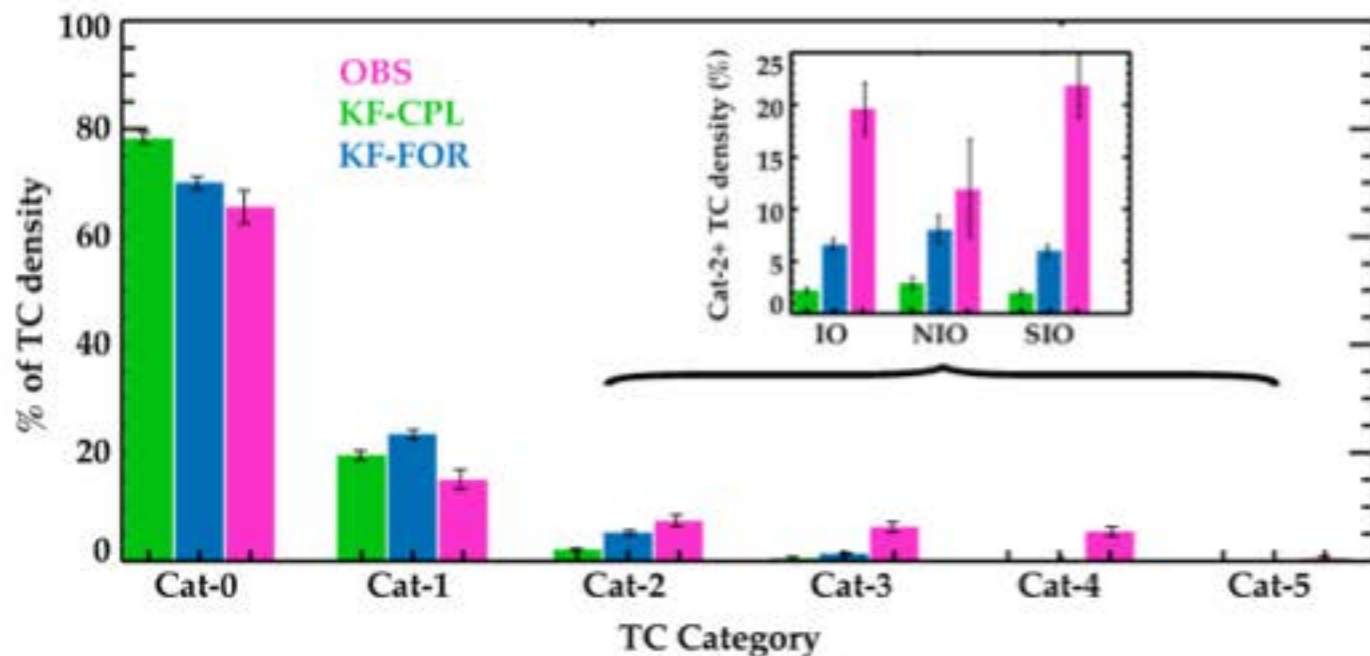
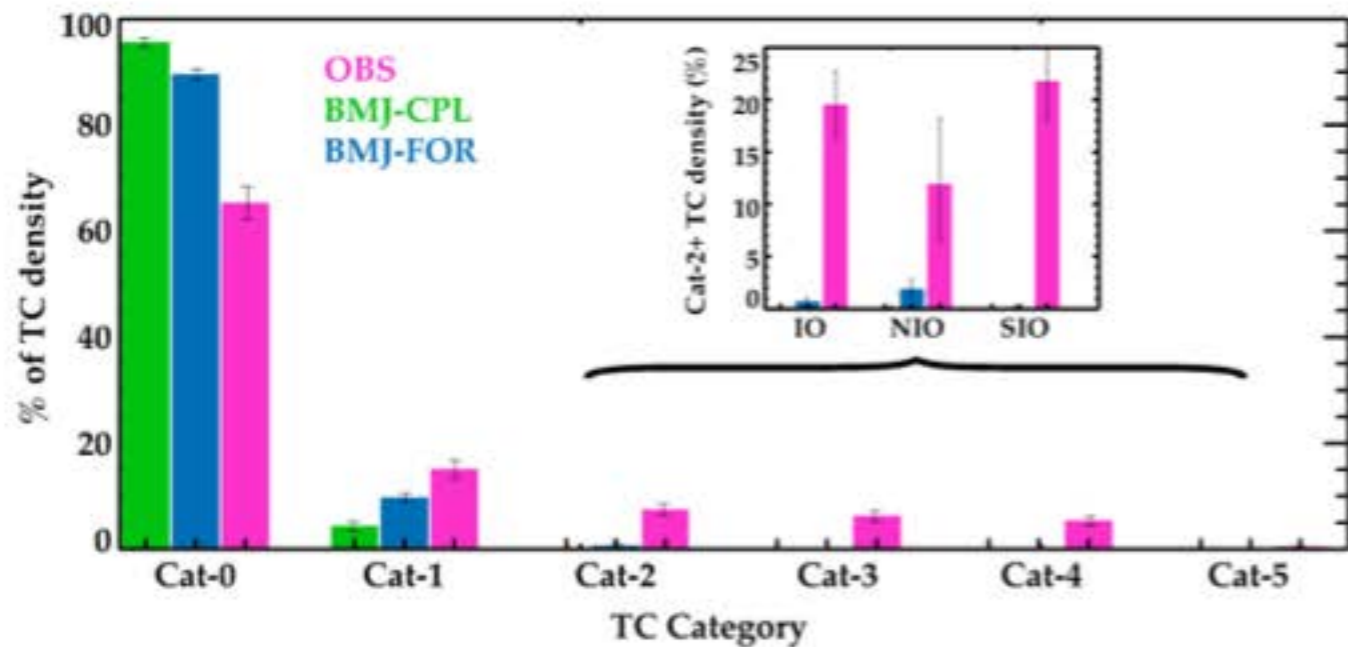


Fig. 8 Distribution of TC central pressure (in number of TC days per year) for the period 1979–1999 and from SPEArTC observations (cyan), the coupled model (blue) and the forced model (green). The 5–95 percentiles bootstrap error bars are also presented

Figure 5: Histogram of the percentage of Indian Ocean TCs as a function of TC intensity, based on maximum TC wind for observations (purple), KF-CPL (green) and KF-FOR (blue) simulations. The inset indicates the percentage of intense TCs (defined as category 2 or above). The whiskers display the 90% confidence interval, computed using a Student's t-test.

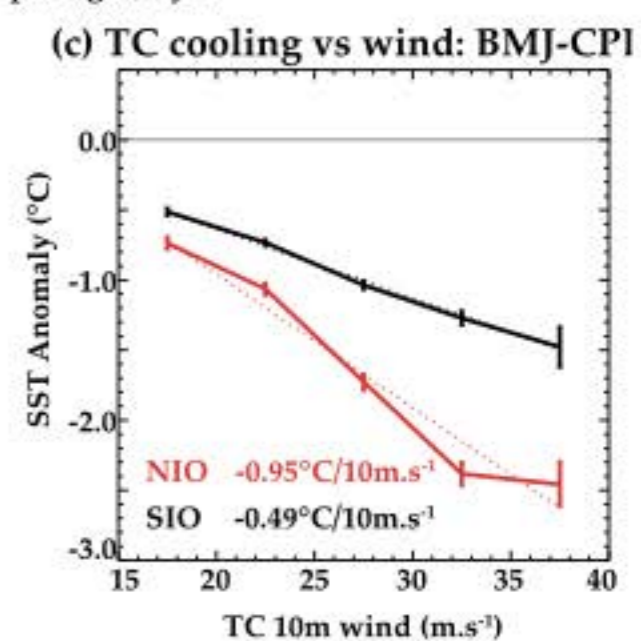
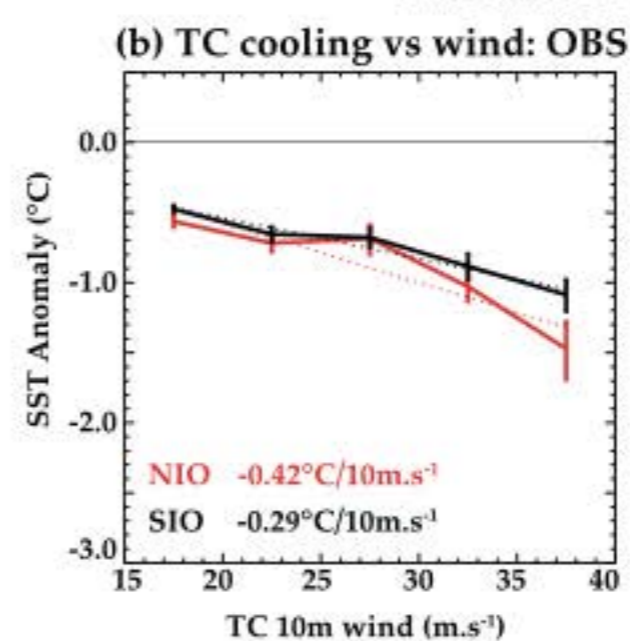
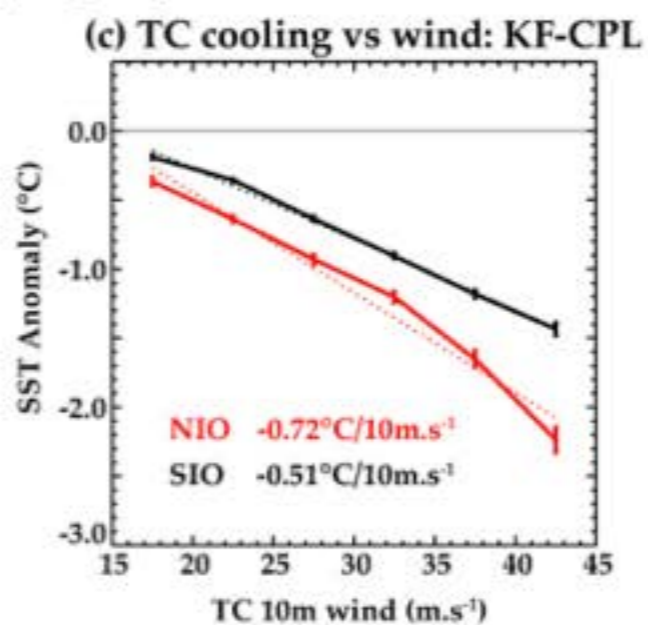
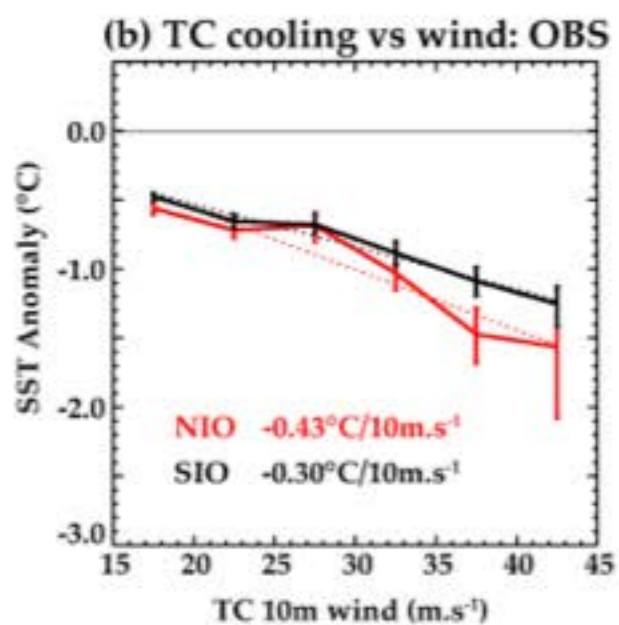
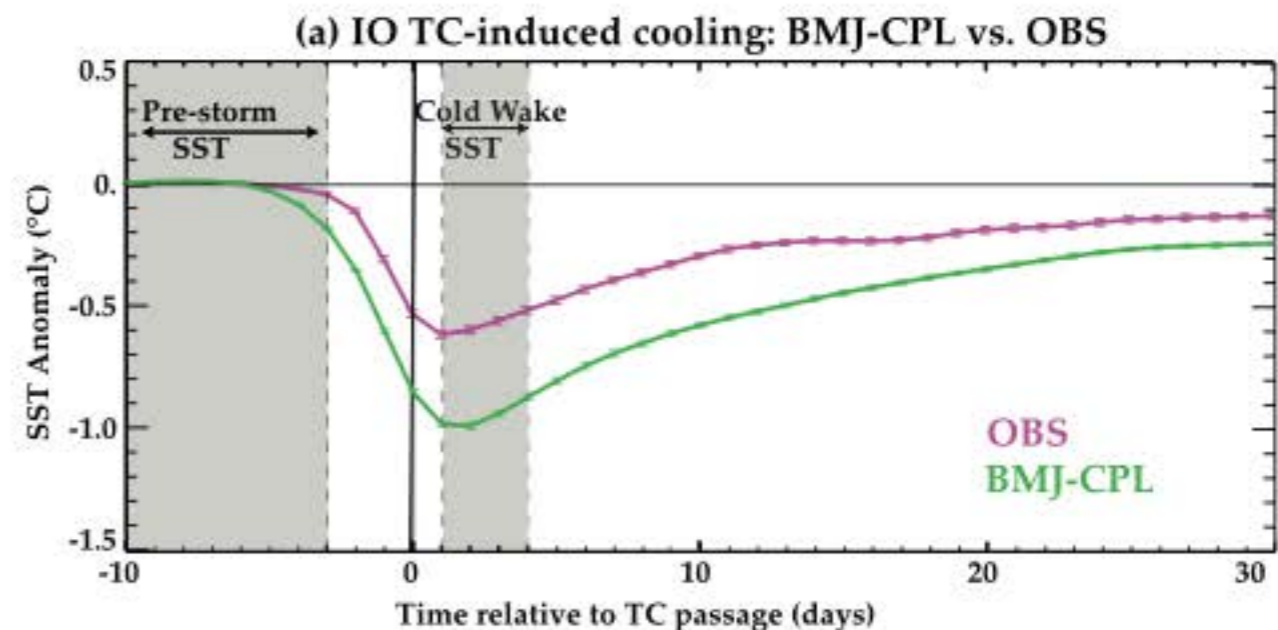
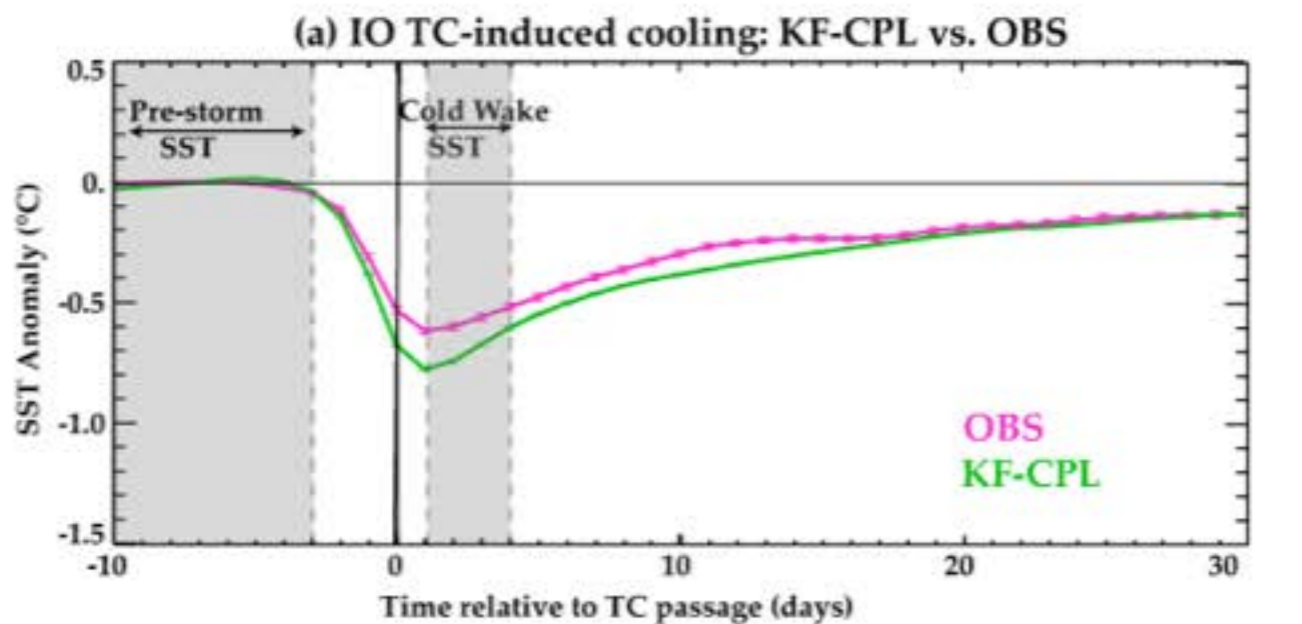


Figure 6: (a) Composite evolution of the TC-induced SST cooling within 200 km of all TC-tracks in the IO (in $^{\circ}\text{C}$) for observations (purple) and the KF-CPL experiment (green). Northern (red) and southern (black) IO mean TC-induced cooling as a function of maximum wind speed for (b) observations and (c) KF-CPL. Whiskers indicate the 90% confidence level from a bootstrap method. The slope of the linear fit is also indicated as a dotted line.

TCs composite in the northern IO

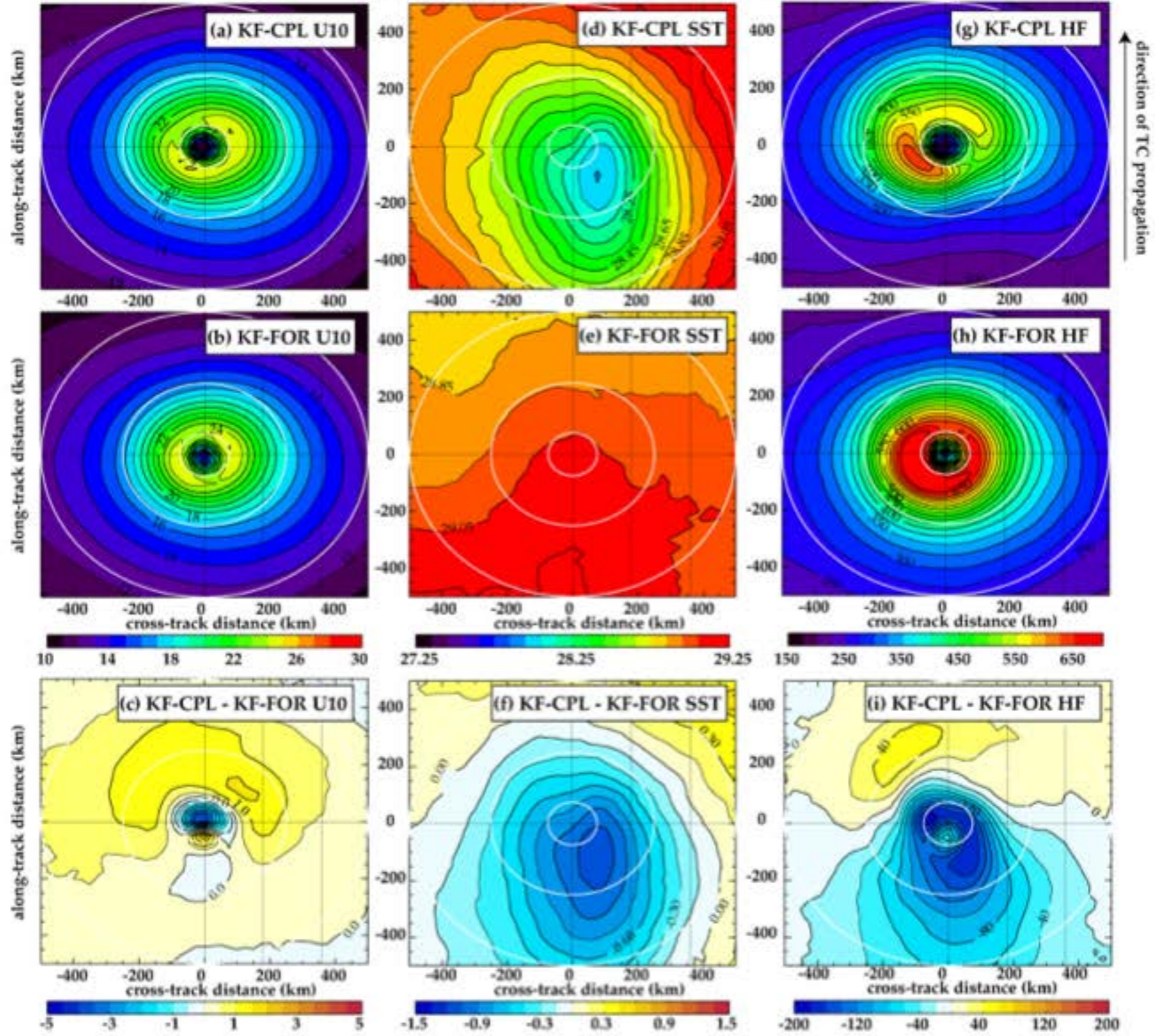
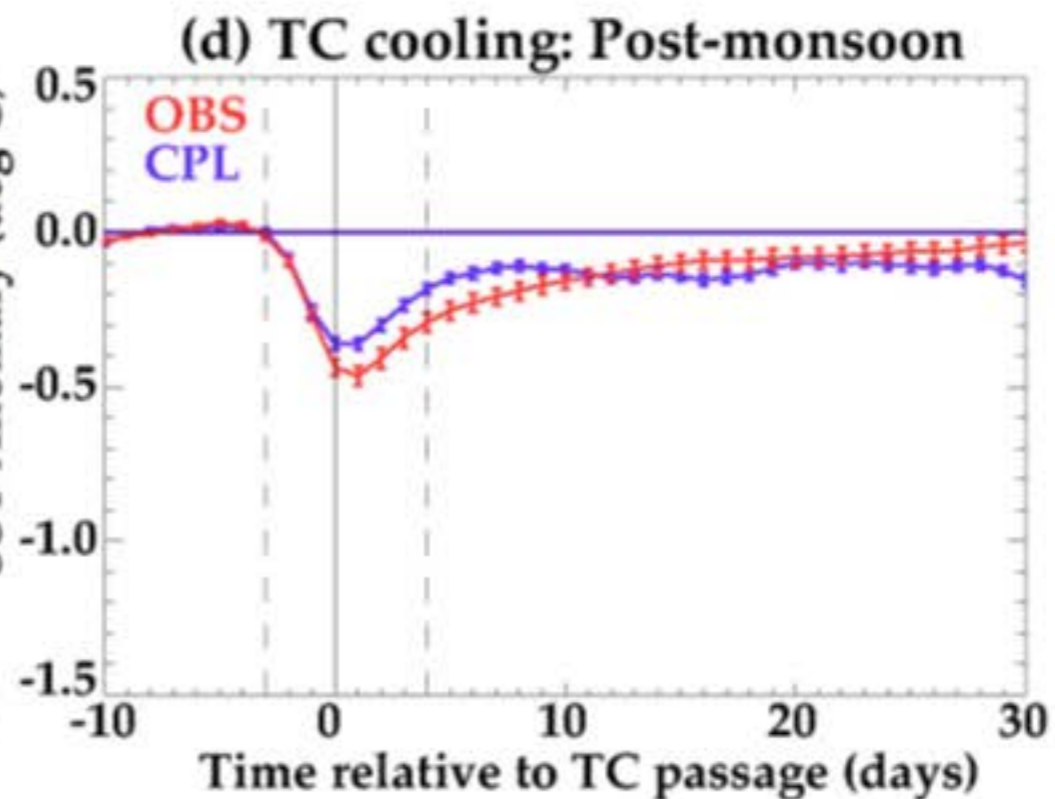
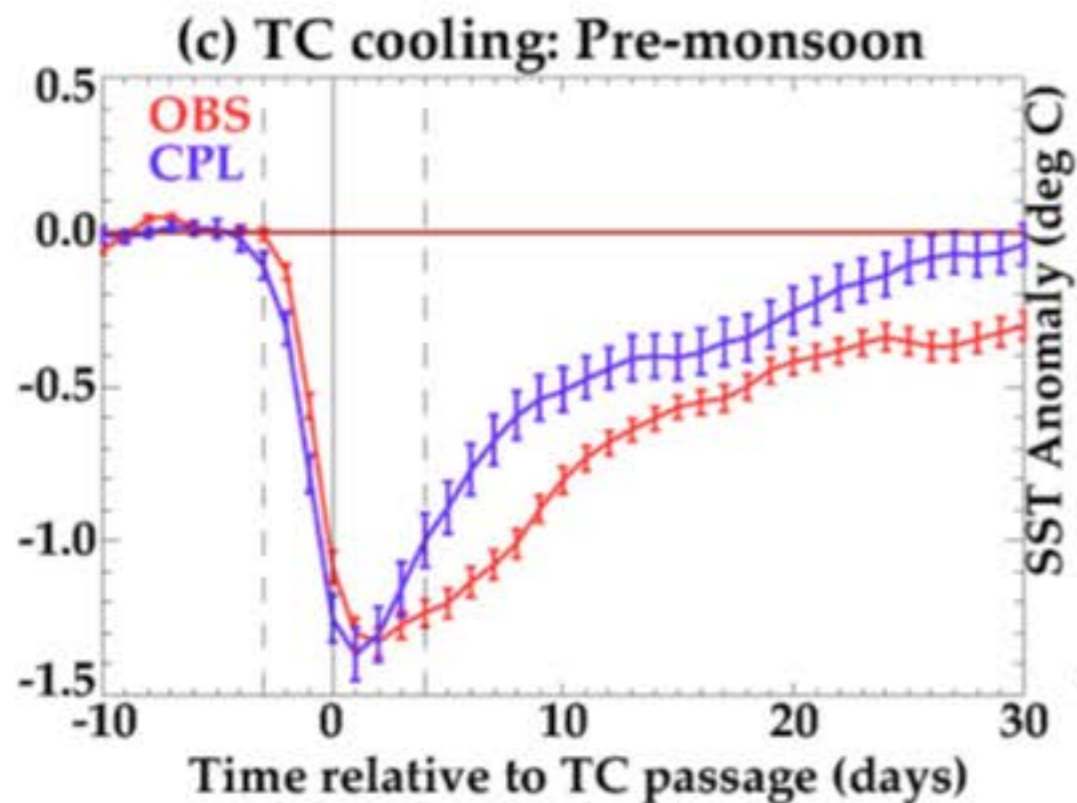
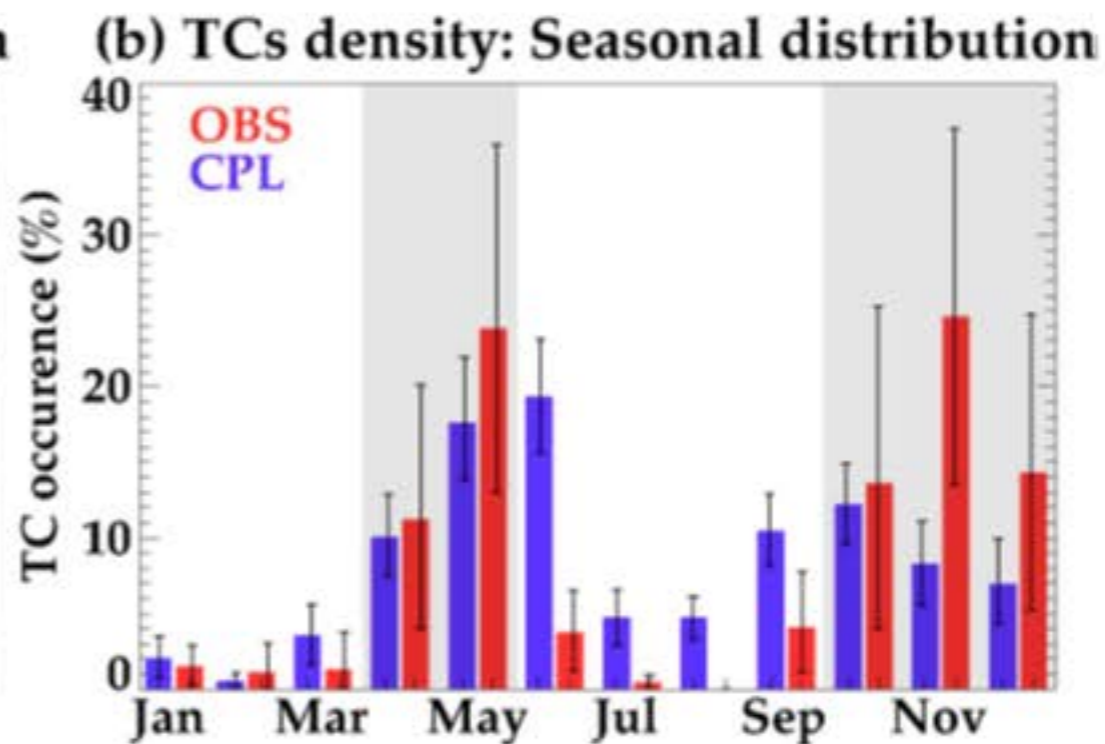
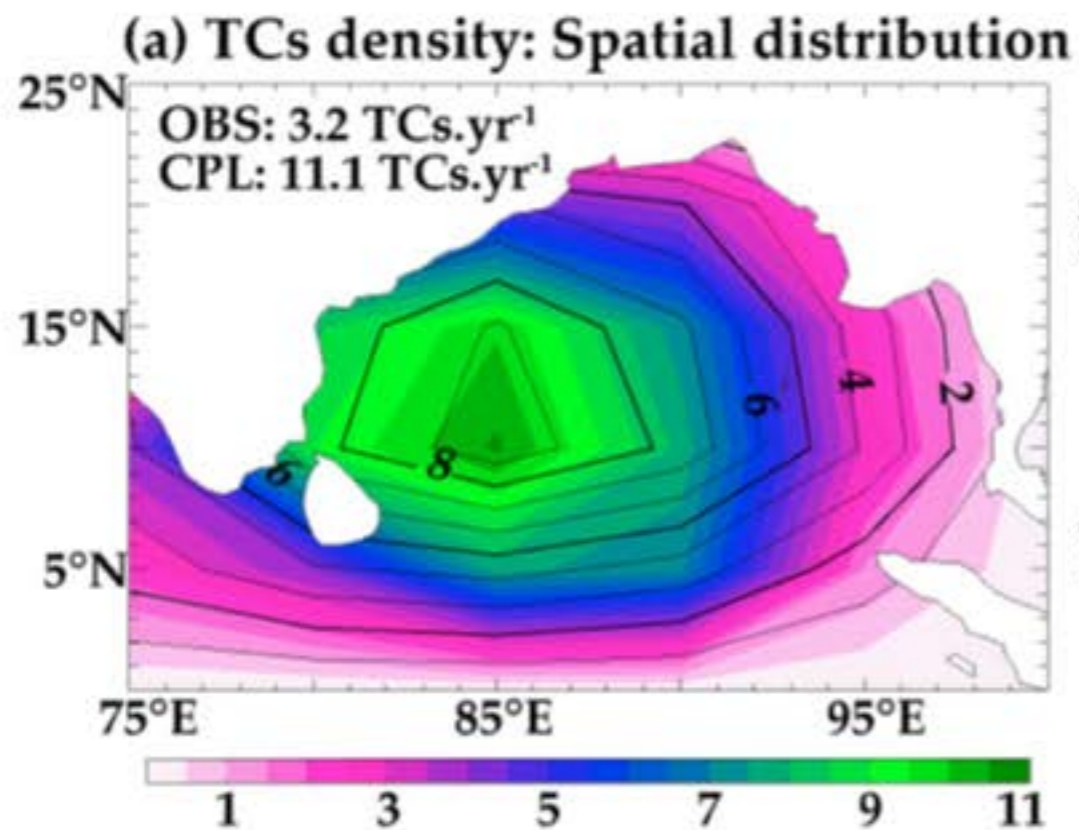
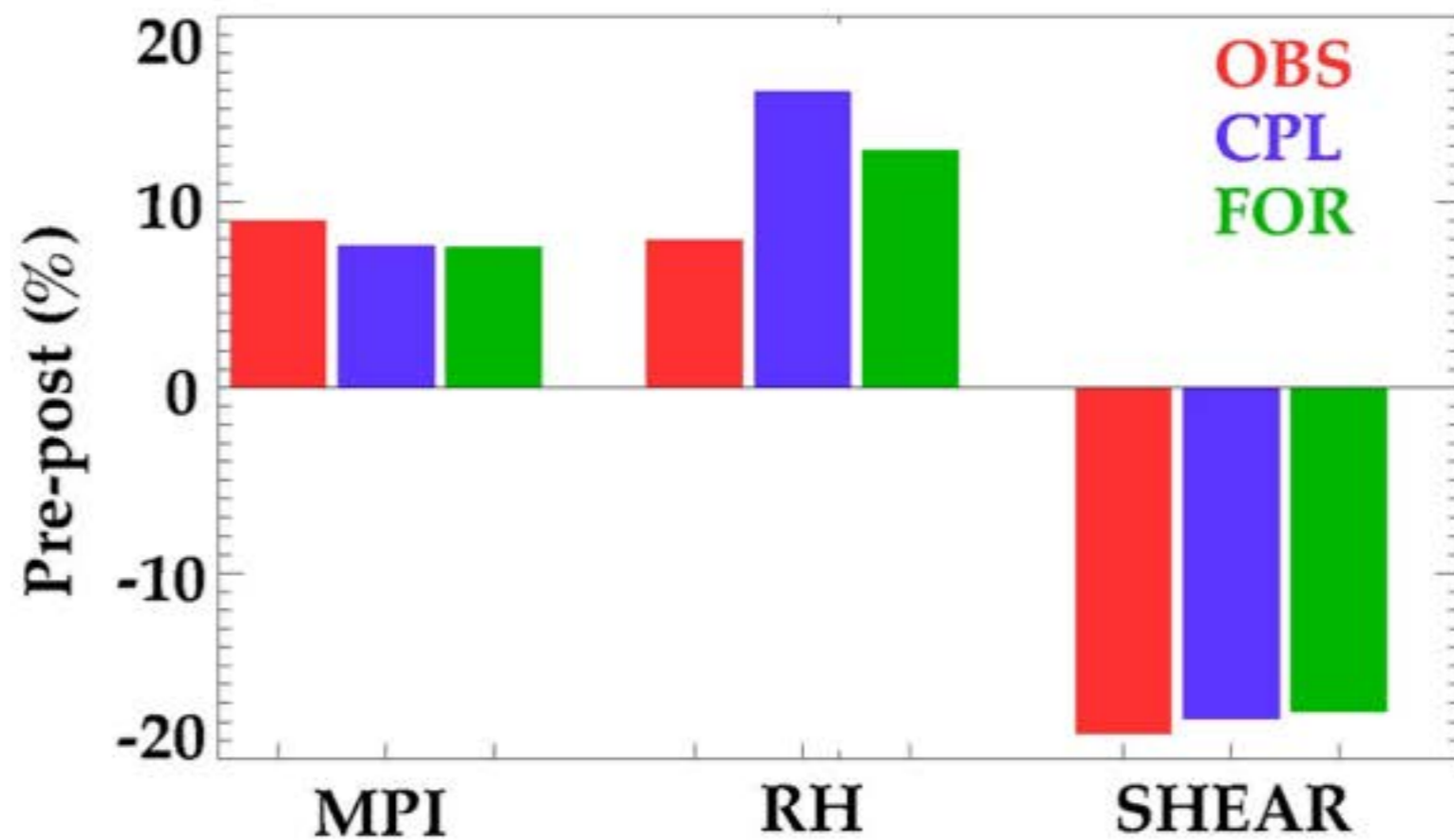


Figure 9: Same as Figure 8 for the northern IO.

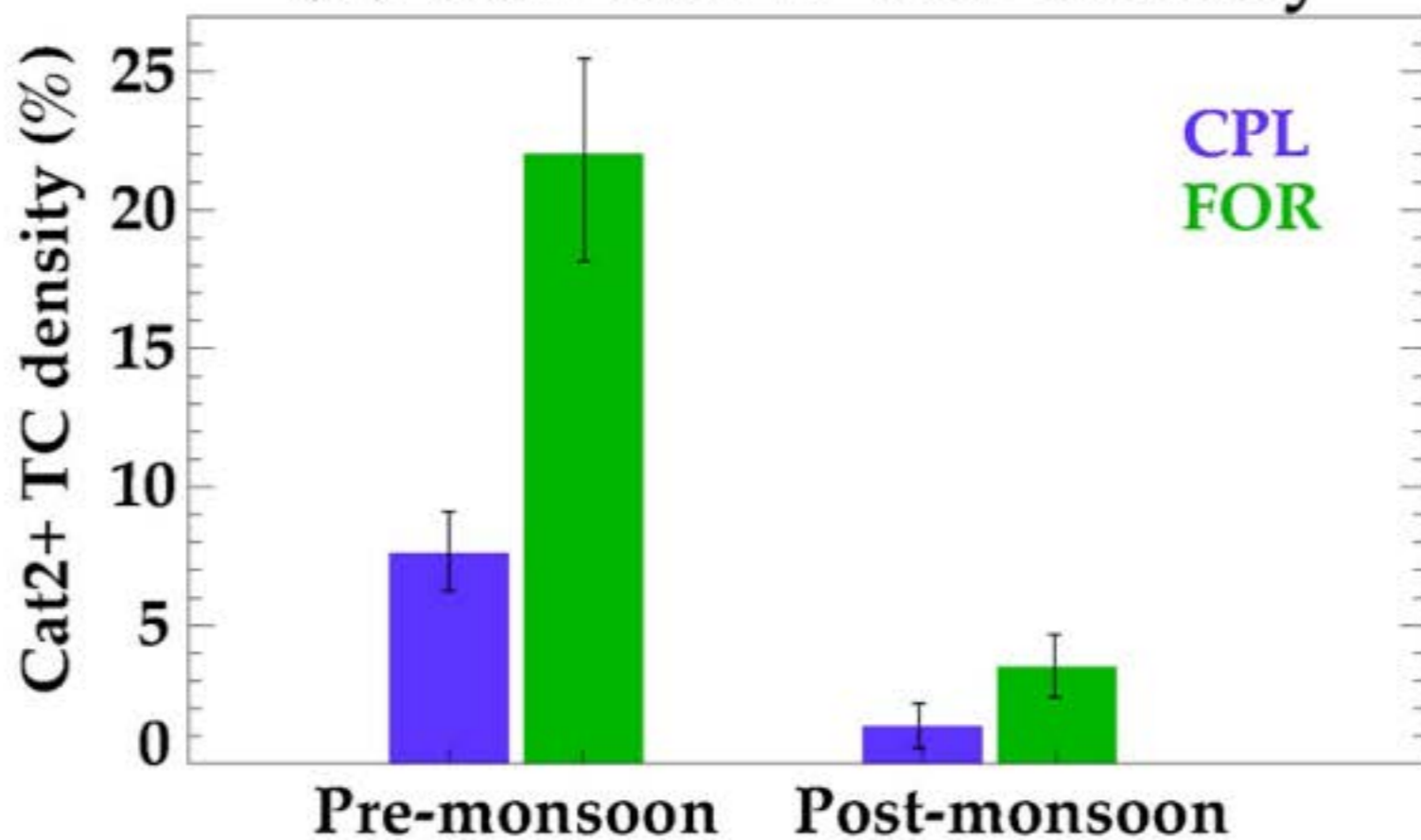
Impact oceanic
stratification



(a) BoB environmental variables



(b) BoB Cat2+ TCs density



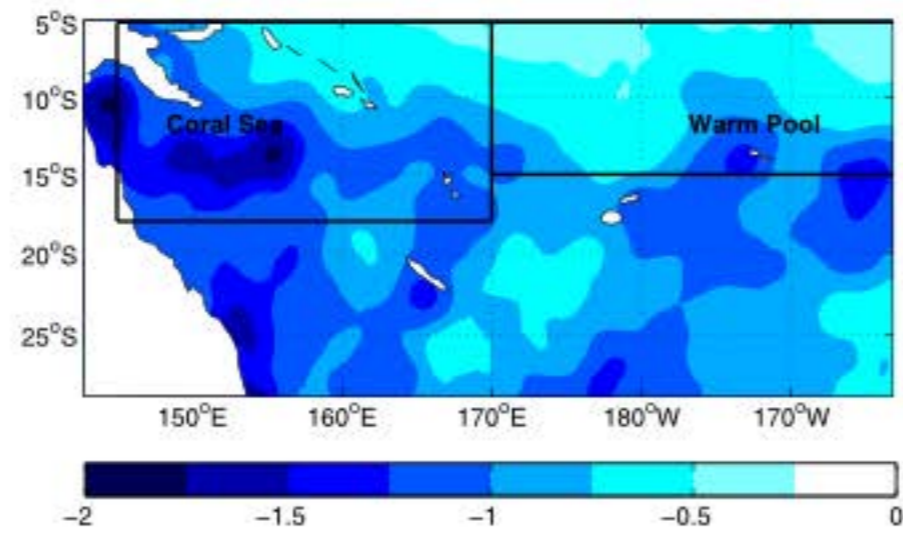
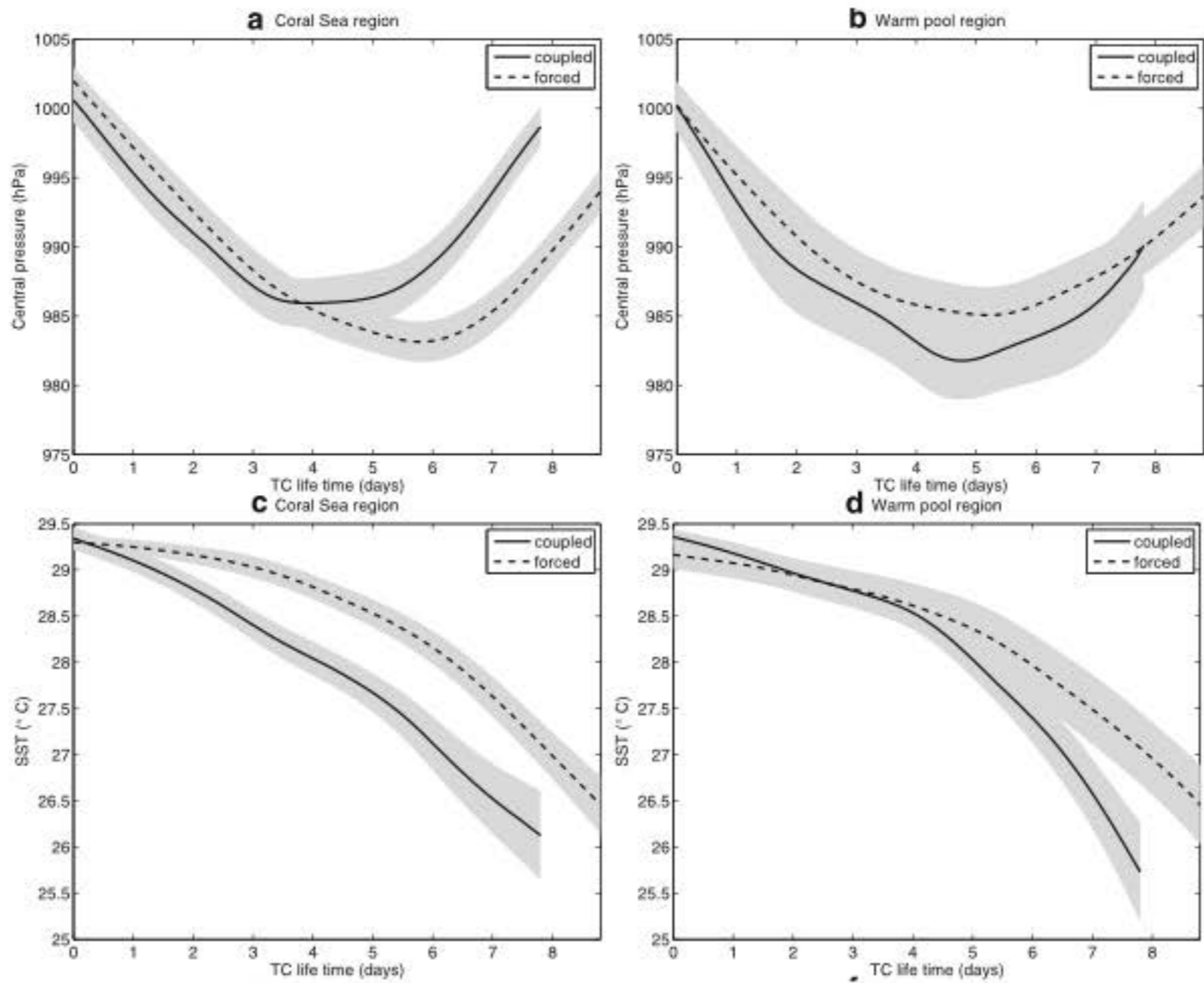


Fig. 10 Composite map of storm-induced SST cooling ($^{\circ}\text{C}$). In the construction of this map, if two TC tracks overlap, the coolings are averaged



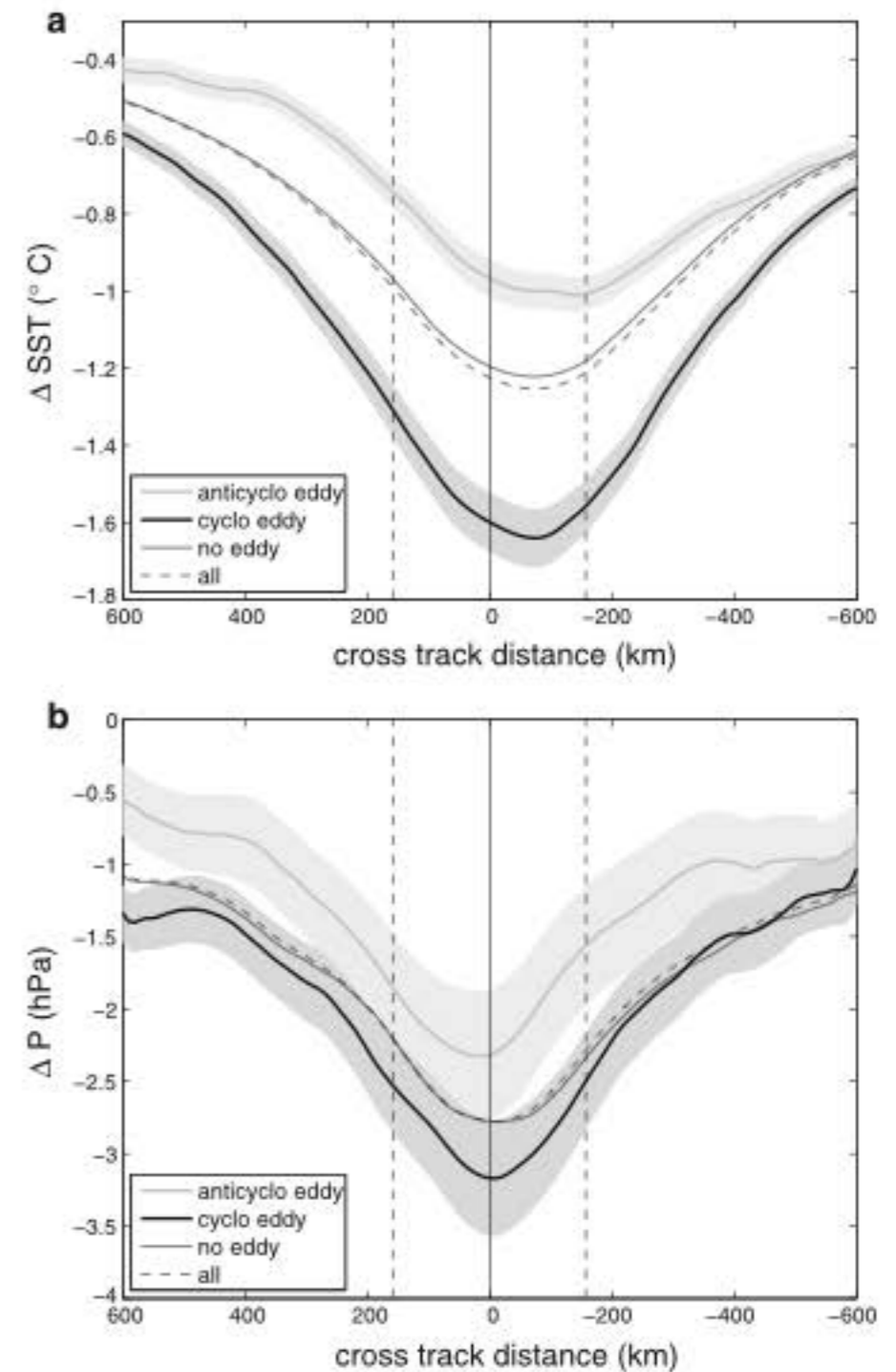
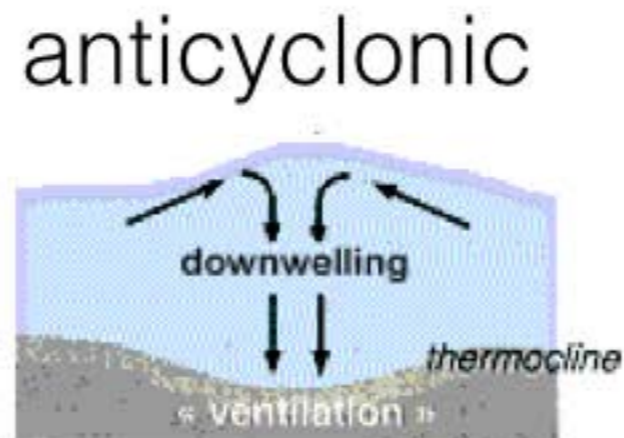
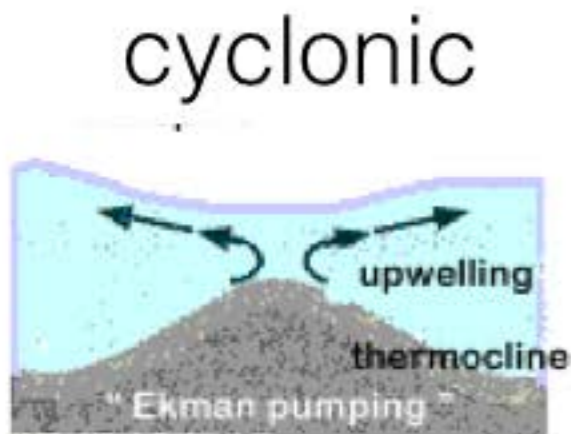
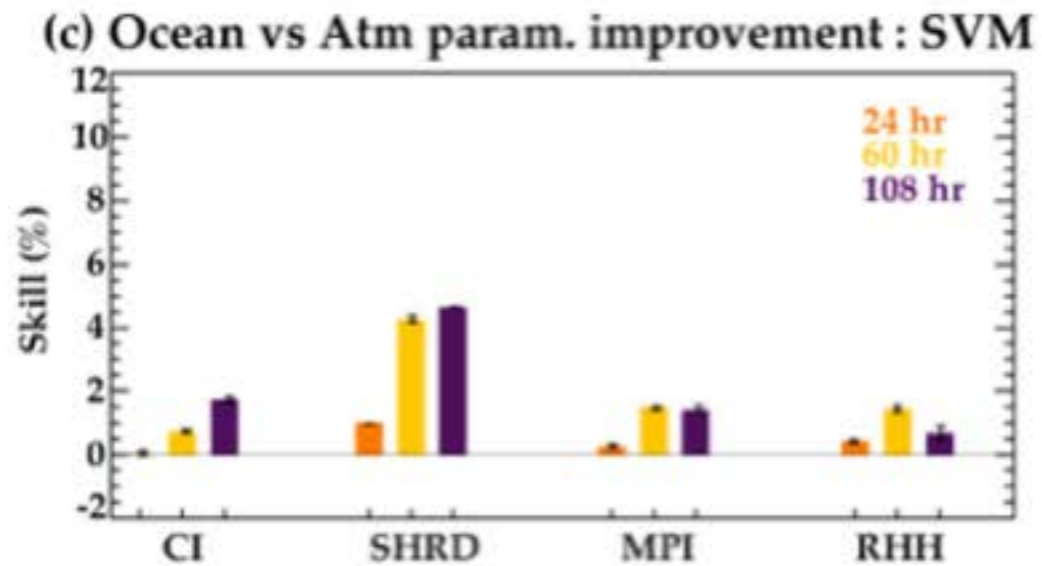
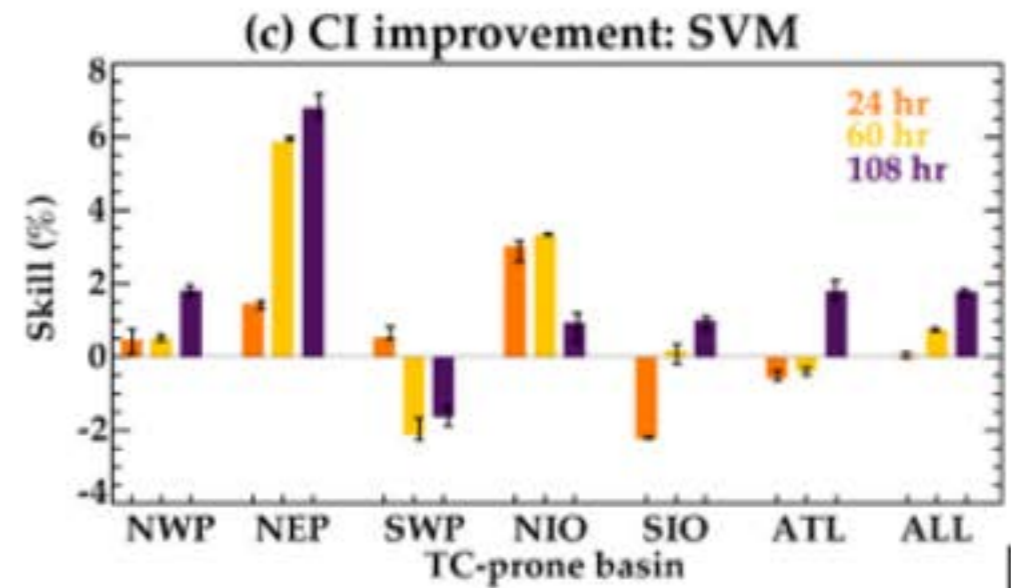
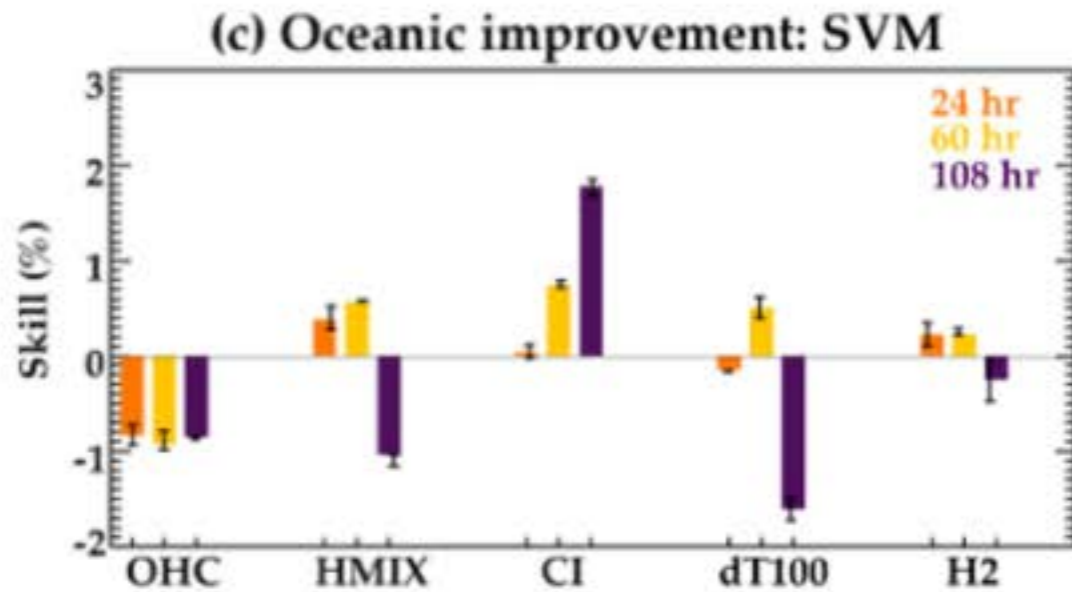


Fig. 13 Cross track composites of **a** SST cooling ($^{\circ}\text{C}$) in the coupled model and **b** TC central pressure differences (hPa) between forced and coupled models. The results are given for all TCs (*dashed line*); TCs located over a cyclonic eddy (*bold black solid line*); TCs located over an anticyclonic eddy (*bold gray solid line*); TCs located over a neutral ocean (neither a cyclonic nor anticyclonic eddy; *thin black solid line*). The standard error of the mean (at 90 % confidence) is presented in *shades of grey*

Air-sea coupling used in Statistical models...



conclusion

- Sensitivity of the results to parameterizations... consequences for our work???
- clear and significant impact of coupling mainly on TC intensity
- Sensitivity to ocean stratification (seasonal variations, regional variations, eddies impact...)